

Oligocene-Miocene biostratigraphy, magnetostratigraphy, and isotopic stratigraphy of the western North Atlantic

Kenneth G. Miller

Lamont-Doherty Geological Observatory of Columbia University, Palisades, New York 10964

Marie-Pierre Aubry

Centre de Paléontologie Stratigraphique et Paléoécologie de l'Université Claude Bernard, Villeurbanne, France,
and Woods Hole Oceanographic Institution, Woods Hole, Massachusetts 02543

M. J. Khan*

Lamont-Doherty Geological Observatory of Columbia University, Palisades, New York 10964

A. J. Melillo

Department of Geological Sciences, Rutgers University, New Brunswick, New Jersey 08903

D. V. Kent

Lamont-Doherty Geological Observatory of Columbia University, and Department of Geological Sciences, Columbia University,
Palisades, New York 10964

W. A. Berggren

Woods Hole Oceanographic Institution, Woods Hole, Massachusetts 02543,
and Department of Geological Sciences, Brown University, Providence, Rhode Island 02912

ABSTRACT

Magnetostratigraphic records from western North Atlantic Deep Sea Drilling Project (DSDP) Sites 563 and 558 are correlated with the geomagnetic polarity time scale (GPTS; Berggren et al., 1984a, 1984b) using marine magnetic anomalies and selected biostratigraphic datum levels. The magnetostratigraphy established is used to make direct magnetobiostratigraphic correlations that agree with previous Oligocene-early Miocene studies. However, we show that Zones NN8 partim and NN9 and associated Epoch 11 correlate with Magnetic Anomaly 5 (= Chron C5n). This contrasts with previous indirect correlations of Epoch 11 with Anomaly 5A and requires an upward adjustment of 1.5–2.0 m.y. for middle-late Miocene calcareous nannofossil zones. We correlate the middle/late Miocene boundary with Zone NN8 and earliest Chron C5n (10.4 Ma).

INTRODUCTION

Integrated magnetobiostratigraphic studies of marine sediments (Lowrie et al., 1982; Berggren et al., 1983; Poore et al., 1983; Shackleton et al., 1984a) have allowed the direct correlation of calcareous microfossil first (FO) and last occurrences (LO) with the GPTS. This provides numerical age estimates for biozonal boundaries and chronostratigraphic units by biostratigraphic correlation with the stratotypes. Direct correlations of microfossils with the GPTS have yielded relatively precise numerical age estimates of Oligocene-lower Miocene stage boundaries. However, previous middle-upper Miocene calcareous magnetostratigraphic correlations have been indirect (Ryan et al., 1974; Theyer and Hammond, 1974), resulting in equivocal correlations of the stratotype Langhian, Serravalian, and Tortonian Stages with the GPTS (Berggren et al., 1985a, 1985b). DSDP Sites 563 (lat 33°39'N, long 43°46'W; 3796

m present water depth) and 558 (lat 37°46'N, long 37°21'W; 3754 m present water depth) provide relatively continuous North Atlantic Oligocene-Miocene records that yield direct correlations of nannofossils, planktonic foraminifera, and isotopes with the GPTS.

METHODS

Details of magnetostratigraphic techniques and interpretations are given by Khan et al. (1985). Normal and reversed polarity magnetozones are interpreted from remanent inclination directions after selective demagnetization in alternating fields of 10 mT or higher (Figs. 1, 2). Both holes 563 and 558 were rotary cored. Although the effects of drilling disturbance and coring gaps are difficult to avoid completely, we constructed a consistent magnetostratigraphic framework and correlated it with magnetochrons of the GPTS; chron terms and GPTS are derived from Berggren et al. (1985a, 1985b). Planktonic foraminiferal studies (Miller, Melillo, Berggren) and nannofossil studies (Aubry) are based upon qualitative analyses (samples indicated in Figs. 1, 2). Oxygen and carbon isotope analyses were performed upon the benthic foraminifera *Cibicidoides* (Figs. 1, 2). Procedures and data are given by Miller and Fairbanks (1985). Possible recrystallization of foraminifera was noted in the middle Miocene and lower Oligocene at Site 558 (Fig. 2); therefore, interpretations of Site 558 are tentative. Both sites were located at abyssal depths (>2 km) during the Oligocene-Miocene.

MAGNETOCHRONOLOGIC FRAMEWORK

In general, Site 558 contains the better Oligocene magnetostratigraphic record, whereas Site 563 contains the better Miocene record. Apparently, both sites have lower Miocene unconformities.

Site 558 was drilled on crust located between Magnetic Anomalies 13 and 15 (Bougault, Cande, et al., 1985), suggesting that the normal magnetozones found near the base of the hole (Fig. 2) is Chron C13n. With this constraint, the observed polarity sequence can be correlated with the GPTS. Several biostratigraphic datum levels, which have been

*Present address: Peshawar University, Peshawar, Pakistan.

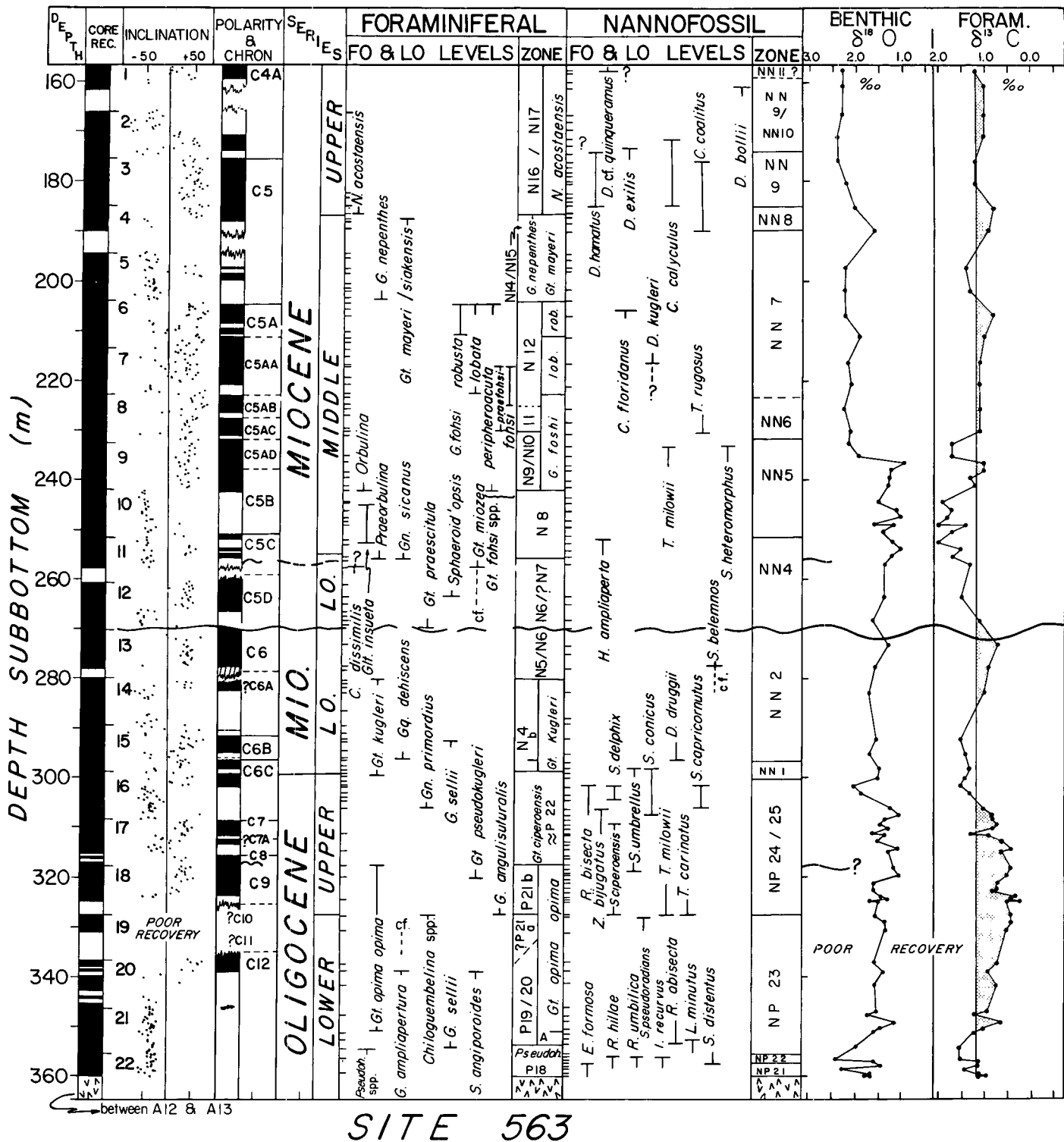
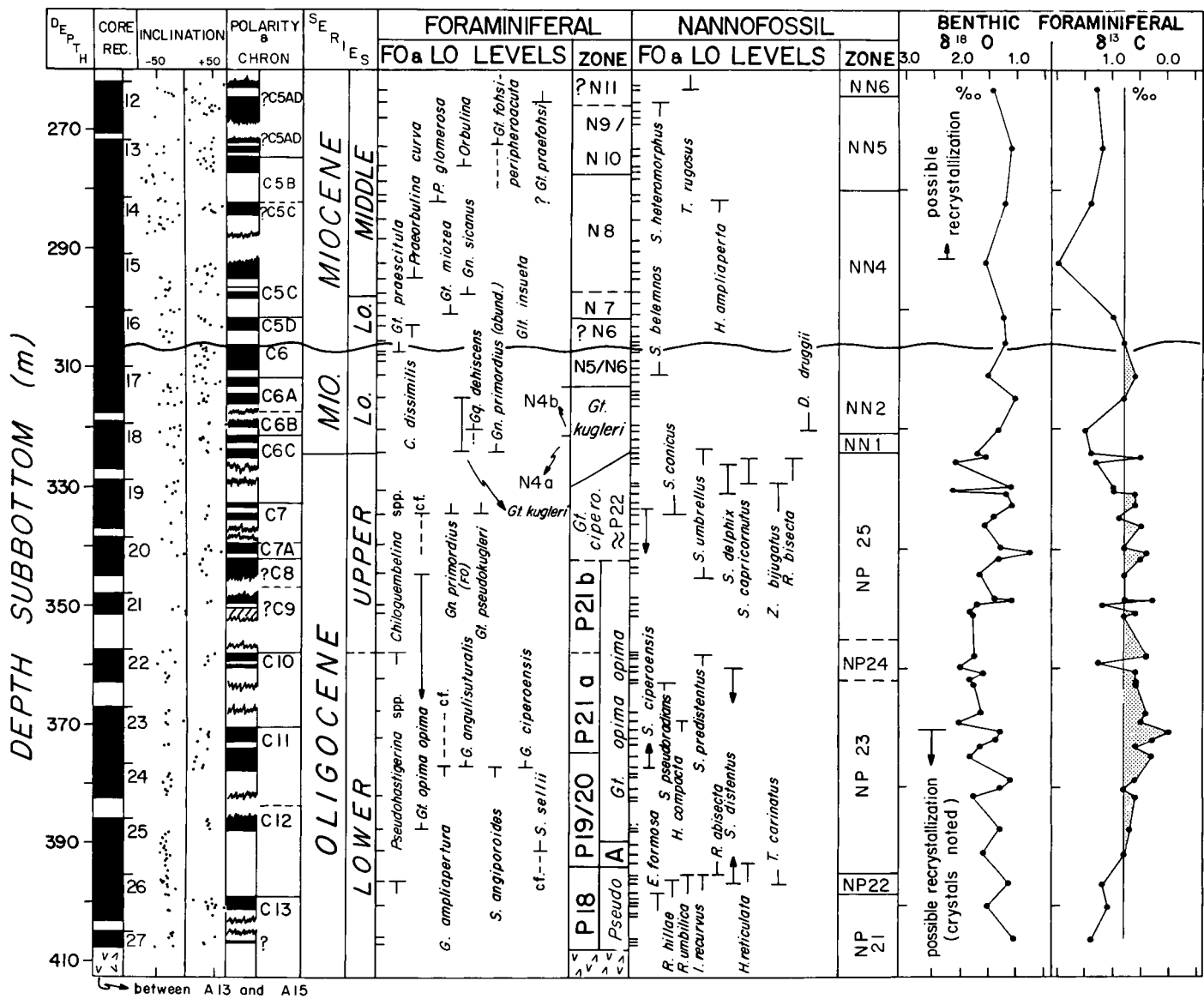


Figure 1. Lower Oligocene–upper Miocene magnetostratigraphy, biostratigraphy, and isotopic stratigraphy of western North Atlantic Site 563. Chron terms after Berggren et al. (1985a, 1985b): e.g., Chron C5 is a time unit corresponding to Anomaly 5 time. Shaded areas in CORE REC. column indicate recovery. Wavy lines at 270 msb, 255 msb, and 317 msb indicate unconformities. Positive or negative inclination values indicate normal (shaded) or reversed (open) magnetic polarity; slashes indicate uncertain polarity. Magnetostratigraphic data after Khan et al. (1985); correlations with Chrons C5AD through C5AA are tentative. Isotopic data relative to PDB standard after Miller and Fairbanks (1985); mean value drawn through carbon isotope data; stipple = values less than mean. Enriched oxygen isotope interval noted near lower/upper Oligocene boundary at Site 558 is missing from this site because of poor recovery (Chrons C10 and C11 are missing). Tick marks indicate levels of samples examined for biostratigraphy. Nannofossil zonation scheme after Martini (1971); planktonic foraminiferal scheme modified after Bolli (1957), Stainforth et al. (1975), Blow (1979), and Kennett and Srinivasan (1983). A = *Globigerina ampliapertura* Zone. FO = first occurrence of taxa at this location. LO = last occurrence. Series boundaries drawn using magnetobiostratigraphic criteria of Berggren et al. (1985a, 1985b) and criteria given in text.

correlated firmly with the GPTS elsewhere, support our correlations. The LO of *Chiloguembelina* spp. at Sites 558 (Fig. 2), 516 (Pujol, 1983), and 522 (Poore et al., 1983) correlates with Chron C10n. The FO of *Globorotalia kugleri* was in Chron C6Cn at Sites 558 (Fig. 2), 563 (Fig. 1), and 516 (Berggren et al., 1983). Nannofossils indicate the absence of Zone NN3, suggesting a hiatus for Chron C5E (about 1 m.y.; Figs. 1, 2). The FO of *Orbulina* spp. was in Chron C5Bn, in agreement with the correlations at Sites 563 (Fig. 1) and 522 (Poore et al., 1983). Poor recovery makes identification of Chrons C5AD-C5A problematic (Fig. 2); a long normal magnetozone occurring in cores 4-6 is correlated with Chron C5n.

Site 563 was drilled between Anomalies 12 and 13 (Bougault, Cande, et al., 1985), suggesting that the long reversed magnetozone noted below 339 m subbottom (msb) represents Chron C12r (Fig. 1). Poor core recovery precludes identification of Chrons C10 and C11, and

Chron C9 and C8 are apparently concatenated (Fig. 1). Correlations near the lower/upper Oligocene boundary are therefore uncertain. The FO of *Globorotalia kugleri* is associated with the lower of two normal magnetozones correlated with Chron C6C, agreeing with correlations noted above. As at Site 558, a missing section or sections is recognized in the lower Miocene (between cores 12 and 13 and/or near the base of core 11). The absence of nannofossil Zone NN3 suggests that the unconformity occurs at the base of core 12 with Chron C5E (about 1 m.y.) missing (Fig. 1). The first occurrence of *Orbulina* spp. occurs at the base of a normal magnetozone correlated with Chron C5Bn, as noted at Sites 521 and 558. The predominantly normal interval punctuated by short reversed magnetozones in cores 6 through 9 is tentatively correlated with Chrons C5AD through C5A. Therefore, the long normal magnetozone encountered in cores 3 and 4 (Fig. 1) is correlated with Chron C5n (= Anomaly 5).



HOLE 558

Figure 2. Lower Oligocene-middle Miocene stratigraphy of western North Atlantic Site 558. Description as in Figure 1. Wavy line at 307 m indicates unconformity. Recovery of cores 7-11 was poor; therefore, data for cores 1-11 are not shown but are available on request from Miller.

PLANKTONIC FORAMINIFERAL BIOSTRATIGRAPHY

The classic tropical zonation of Bolli (1957) and Blow (1979) are only partially applicable at these mid-latitude, subtropical gyre locations. *Globorotalia opima opima* (>390 μm) first occurred in Chron C12n at Site 558 (Fig. 2). At Site 563 (Fig. 1) it apparently occurred even earlier (Chron C12r) immediately after the LO of *Pseudohastigerina* spp. (effectively eliminating *G. ampliapertura* Zone). *Globigerinoides primordius* first occurred in late Oligocene Chron C6Cr at both sites (Figs. 1, 2), in agreement with Berggren et al. (1985a, 1985b and references therein), but in contrast to Blow (1979). The FO of *Gt. fohsi periphroacuta* and the FO *Orbulina* spp. are closely associated at both sites, suggesting the elimination of Zone N9. The FO of *Globigerina nepenthes* immediately after both Chron C5An and the LO of *Globorotalia fohsi* ssp. (Fig. 1) suggest the elimination of Zone N13 (sensu Kennett and Srinivasan, 1983). Similarly, Zone N15 (sensu Blow, 1979) is eliminated by the virtual overlap of *Gt. mayeri/siakensis* and *Neogloboquadrina acostaensis* (Fig. 1). As the magnetostratigraphic record is fairly complete, we believe that the elimination of these zones at these locations is due not to hiatuses but to latitudinal diachrony. Further documentation of these apparently diachronous ranges is needed.

At Site 563, *Catapsydrax dissimilis* disappears just below the FO of *Praeorbulina* spp. This indicates that Zone N7 is missing; it is apparently present at Site 558. This suggests a hiatus (Chron C5C partim) at Site 563.

Controversy surrounds the placement of the Oligocene/Miocene boundary. Berggren et al. (1985a, 1985b) recognized the boundary using the FO of *Globorotalia kugleri*, whereas Kennett and Srinivasan (1983) used the FO of *Globoquadrina dehiscens*. We note that at Sites 563 and 558 (Figs. 1, 2) these events were linked with Chrons C6Cn (at 23.7 Ma) and C6Br (at 23.2 Ma), respectively, and that both events postdated a sharp oxygen isotope enrichment that occurred in Chron C6Cr (at 24–25 Ma; Figs. 1, 2). These stratigraphic relationships show a remarkable consistency between sites. We arbitrarily place the boundary in Chron C6Cn in association with the FO of *G. kugleri*.

NANNOFOSSIL BIOSTRATIGRAPHY

We were able to apply Martini's (1971) standard zonation at Sites 558 and 563 except that (1) in the lower–middle Miocene poor preservation precluded accurate recognition of the FO and LO of some stratigraphically useful discoasters (i.e., *D. formosus*, *D. exilis*, *D. variabilis*, *D. challengerii*); (2) in the upper Miocene preservation is poor; (3) some local early FO were observed (e.g., *Sphenolithus ciperoensis*, Site 558; Fig. 2); and (4) some index species were rare (*S. distentus*, *Helicosphaera ampliapertura*, *D. kugleri*). On the basis of nannofossil evidence, a lower Miocene unconformity is postulated for both sites (Zone NN3 is missing). At Site 558 (Fig. 2), reworking is suggested above the unconformity (core 16-4) by the anomalous co-occurrence of *Triquetrorhabdulus carinatus*, *S. belemnus*, and *S. heteromorphus*. At Site 563, Zone NN2 (with *T. carinatus* and *S. belemnus*) is overlain by Zone NN4 (with *S. heteromorphus* and *H. ampliapertura*).

Diverse, well-preserved Oligocene and lower–upper Miocene nannofossils allow firm correlations. The LO of *Reticulofenestra bisecta* is closely associated with the FO of *Globorotalia kugleri* and Chron C6Cn as noted at Sites 516 and 522 (Berggren et al., 1983; Poore et al., 1983; see above). At Site 563, Zones NN8 partim and NN9 are associated with Chron C5n (Fig. 1); this is supported by correlations at Sites 521A, 519 (Hsü, La Brecque, et al., 1984), and 558. Our correlations suggest the following revised magnetostratigraphic boundary ages: (1) NN7/NN8 = 10.8 Ma, and (2) NN8/NN9 = 10.0 Ma.

ISOTOPIC STRATIGRAPHY

At site 563, high $\delta^{18}\text{O}$ values (>2.0‰) in benthic foraminifera occur in the lower Oligocene (within Chron C12r), upper Oligocene

(Chron C6Cr; magnetostratigraphic age estimate 25–24 Ma), and middle Miocene (Chron C5Bn; age estimate 14.7 Ma) to the top of the hole (Fig. 1). At Site 558, the upper Oligocene (25–24 Ma) enriched interval also is present; in addition, an enriched interval occurs across the lower /upper Oligocene boundary (Chron C11n–C9n; 31–28 Ma; Fig. 2). This latter interval correlates with an enriched interval noted at equatorial Pacific sites (Keigwin and Keller, 1985; Miller and Thomas, 1985). The earliest Oligocene enriched interval may correlate with high $\delta^{18}\text{O}$ values noted elsewhere following the Eocene/Oligocene boundary increase (Keigwin, 1980; Poore and Matthews, 1984; Snyder et al., 1984, among others), although the highest values elsewhere occurred at ca. 36–35 Ma, and a precise correlation is not possible. The 14.7 Ma enrichment correlates precisely with the well-established middle Miocene ^{18}O enrichment (Savin et al., 1975, 1981; Shackleton and Kennett, 1975, among others). The 25–24 Ma enriched interval has not been noted elsewhere, presumably owing to lack of study of upper Oligocene sediments. The intervals of enriched ^{18}O are interpreted as periods of continental glaciation and lowered sea level at ca. 36–35, 31–28, 25–24, and younger than 14.7 Ma (Miller and Fairbanks, 1983, 1985; Keigwin and Keller, 1984; Miller and Thomas, 1985; Poore and Matthews, 1984; Shackleton et al., 1984b).

The benthic foraminiferal $\delta^{13}\text{C}$ record documents three cycles of enrichments and depletions, with enriched intervals in Chron C12n (34.5–33 Ma), C6An–C6Cr (25.5–22 Ma), and C5D–C5AD (18–14 Ma) (Figs. 1, 2). Similar $\delta^{13}\text{C}$ fluctuations occurred in benthic and planktonic foraminifera from the Atlantic, Pacific, and Indian oceans, indicating that these cycles represent global changes in the mean ^{13}C of seawater (Miller and Fairbanks, 1985).

DISCUSSION: IMPLICATIONS FOR THE AGE OF THE MIDDLE/LATE MIOCENE BOUNDARY

Our Oligocene–lower Miocene magnetobiostratigraphic correlations agree with previous studies (Lowrie et al., 1982; Berggren et al., 1983; Poore et al., 1983; Pujol, 1983). However, our middle–upper Miocene correlations conflict with those of previous workers. In particular, Ryan et al. (1974) suggested that Epoch 9 was correlated with Anomaly 5 and Epoch 11 with Anomaly 5A. We reevaluate this interpretation based on the following: (1) Zone NN9 is associated with Epoch 11 at piston core RC12-65 (Ryan et al., 1974), and (2) at Sites 563 (Fig. 1) and 558, Zones NN8 (partim) and NN9 are associated with a long normal magnetozone that we correlate with Chron C5n (Anomaly 5). We suggest that Epoch 11 correlates with Anomaly 5 as suggested by Foster and Opdyke (1970). The upward shift of Zone NN9 relative to the GPTS requires an adjustment of 1.5–2.0 m.y. for Zones NN8–NN11.

These magnetobiostratigraphic re-correlations require chronostratigraphic changes. The stratotype Tortonian (lower–upper Miocene) contains Zones NN8 (partim) to NN11 (partim); the base of the stratotype is associated with Zone NN8 and the FO of *N. acostaensis* (references in Berggren et al., 1985a). At Sites 563 and 558, the FO of *N. acostaensis* and Zone NN8 occurred in earliest Chron C5n (Fig. 1). The middle/late Miocene boundary (= basal Tortonian) can be approximated as the earliest Chron C5n; it therefore has a magnetostratigraphic age of 10.4 Ma (GPTS of Berggren et al., 1985a). This contrasts with previous biochronologic age estimates of 11.2 Ma (Berggren and Van Couvering, 1974). Considering the importance of these proposed changes to middle–late Miocene stratigraphy (Berggren et al., 1985a), further studies confirming these magnetostratigraphic relationships are desirable.

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Reviewer's comment

Provides important new data bearing on establishing calibrations between the polarity time scale-biostratigraphic zonations and European stratotypes.

Richard Poore
